

Rainfall Interception in a *Pinus eldarica* Plantation in a Semi-arid Climate Zone: An Application of the Gash Model

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ABSTRACT

Forest canopy rainfall interception loss (I), canopy water storage (S), and the ratio of mean evaporation to mean rainfall intensity (\bar{E}/\bar{R}) are important components of the water balance in arid and semi-arid climate zones. The goal of this project was to quantify I and S and to evaluate the Gash interception model for rainfall interception in a mature semi-arid *Pinus eldarica* Medw afforestation planted in the Chitgar Forest Park near Tehran city, Iran. Measurements of gross precipitation (P_G) and throughfall (TF) were recorded on an event basis from September 2009 to April 2010. For the measurement period, P_G totaled 164.8 mm and I totaled 61.2 mm. I was calculated as the difference between P_G and TF . On the event scale, the ratio of $I:P_G$ ranged between 0.195 and 1, and averaged 0.614. There was a strong logarithmic correlation between $I:P_G$ and P_G ($R^2=0.861$; P value ≤ 0.01). As the size of rainfall events increased, $I:P_G$ decreased. The mean method estimated S to be 1.8 mm. The Gash model accurately estimated I to be within 1.1 mm of the total measured value. The results demonstrate that intercepted rainfall represents a considerable portion of P_G in *P.eldarica* afforested regions of the semi-arid climate zone of Iran where soil moisture is a limiting factor for plant growth and productivity.

Keywords: Afforestation, Canopy water storage, Iran, Mean method, Throughfall.

INTRODUCTION

Afforestation is proposed for arid and semi-arid regions because it is assumed to reduce soil erosion (Zhou *et al.*, 2002), combat desertification (Grünzweig *et al.*, 2003), increase CO₂ fixation (De Los Rios-Carrasco *et al.*, 2009), and provide recreational opportunities (Hüttl *et al.*, 2000). However, afforestation may have undesirable hydrological implications. For example, converting grassland to a forest may reduce groundwater recharge and local water availability because trees extract water from deep layers in the soil profile during drier periods and intercept precipitation

during rainfall events (Iroumé and Huber, 2002).

Rainfall interception loss (I) is the proportion of gross rainfall (P_G) that is intercepted, stored and subsequently evaporated from leaves, branches and stems of vegetation during or following rainfall. In forests, I can be substantial, representing between 10 and 40% of annual P_G (Scatena, 1990; Asdak *et al.*, 1998, Levia and Frost, 2003). Therefore, quantifying the magnitude of I is vital in semi-arid and arid regions where soil moisture is a limiting factor affecting plant growth and productivity (Návar *et al.*, 1999a; 1999b, Carlyle-Moses, 2004).

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Throughfall (TF) is the portion of P_G that directly reaches the forest floor through gaps in the canopy [free throughfall coefficient (p)] (Gash, 1979; Herbst *et al.*, 2008) or water dripping from leaves and branches (canopy drip). Stemflow (SF) is rainfall that reaches the ground by flowing down stems or trunks (e.g. Návar and Bryan, 1990; Návar, 2011). I is estimated using the difference between P_G measured above the canopy and the sum of throughfall (TF) and stemflow (SF) (Hutchinson and Roberts, 1981; Mahendrappa, 1990; Tobón Marin *et al.*, 2000; Xiao *et al.*, 2000; Herbst *et al.*, 2008).

Canopy (water) storage capacity (S) is one of the important variables influencing I (e.g. Rutter *et al.*, 1971; Jackson, 1975; Liu, 1997; Pypker *et al.*, 2005; Crockford and Richardson, 2000). The water stored in the canopy can either evaporate directly to the atmosphere, be absorbed by the canopy, or drop to the ground as throughfall or stemflow. S depends on the characteristics of the intercepting surface. These characteristics include leaf area index (LAI) (Liu, 1998; Llorens and Gallart, 2000; Fleischbein *et al.*, 2005) and leaf shape, bark morphology (e.g. Pypker *et al.*, 2011), dimension and orientation of the branches (Jackson, 1975; Návar and Bryan, 1990), rainfall intensity (Calder *et al.*, 1996; Jackson, 1975) as well as climatic factors such as wind speed (Hörmann *et al.*, 1996; Jackson, 1975).

Various types of regression equations have been proposed to calculate I (Zinke, 1967; Jackson, 1975). However, even for the same vegetation type, the equations frequently differed because of the unique characteristics of each forest stand (van Dijk and Bruijnzeel, 2001). Gash (1979) introduced a simpler storm-based model that has been used with considerable success to estimate I in a wide range of coniferous and broadleaf forests (e.g. Gash *et al.*, 1980; Pearce *et al.*, 1980) as well as tropical rainforests (e.g. Lloyd *et al.*, 1988; Hutjes *et al.*, 1990). However, attempts to use the analytical model in more open forests tended to overestimate I (Teklehaimanot *et al.*, 1991; Gash *et al.*, 1995; Návar, 2012). This led to the development of a 'sparse canopy' version of the Gash model in which evaporation from the wet canopy was considered linearly

dependent on canopy cover fraction (c) (Gash *et al.*, 1995).

Land managers in Iran are faced with year round water shortages because the country has vast expanses of arid and semi-arid regions. During the dry period (May to October), the water shortages become particularly severe. Knowledge about the amount of rainfall intercepted by different tree species will help managers choose suitable species. Furthermore, evidence demonstrating the applicability of the Gash model in this region will provide a tool for land managers to estimate I for different forest types. Currently, *Pinus eldarica* is a widely accepted species for afforestation in Iran, as well as other countries with similar climates (De Los Ríos-Carrasco *et al.*, 2009), because *P. eldarica* tolerates drought (Sardabi, 1998). To our knowledge, a comprehensive investigation on the impact of *P. eldarica* on I has not been reported for forests in Iran, nor in other countries in the region, despite the widespread use of this species in afforestation efforts. Therefore, the objectives of this paper are to (i) quantify how TF and I are partitioned in a planted *P. eldarica* forest located in a semi-arid climate zone of Iran and (ii) determine S using the mean method and (iii) assess the applicability of the Gash model (Gash *et al.*, 1995) for *P. eldarica* forests.

MATERIALS AND METHODS

Site Description

The study occurred in a nearly closed canopied, forty-year-old pure and even-aged *Pinus eldarica* Medw. afforestation located in the Chitgar Forest Park of Tehran, Iran (Figure 1). The pine forest covers 366.5 ha of the Park and represents 45% of the total area. TF measurements were made in a 270 m² plot (35°10' N, 51°10' E, and 1269 m asl). Tree density was 1185 trees ha⁻¹ and the total basal area was 64.5 m² ha⁻¹. Mean tree height and diameter at breast height (DBH) were 11 m and 23.5 cm, respectively. Measurements were performed from September 2009 to April 2010.

From 1996-2010, mean annual precipitation was 267.6 mm (SE: ±20.4 mm) (Chitgar Meteorological Station (35° 44' N, 51° 10' E, and

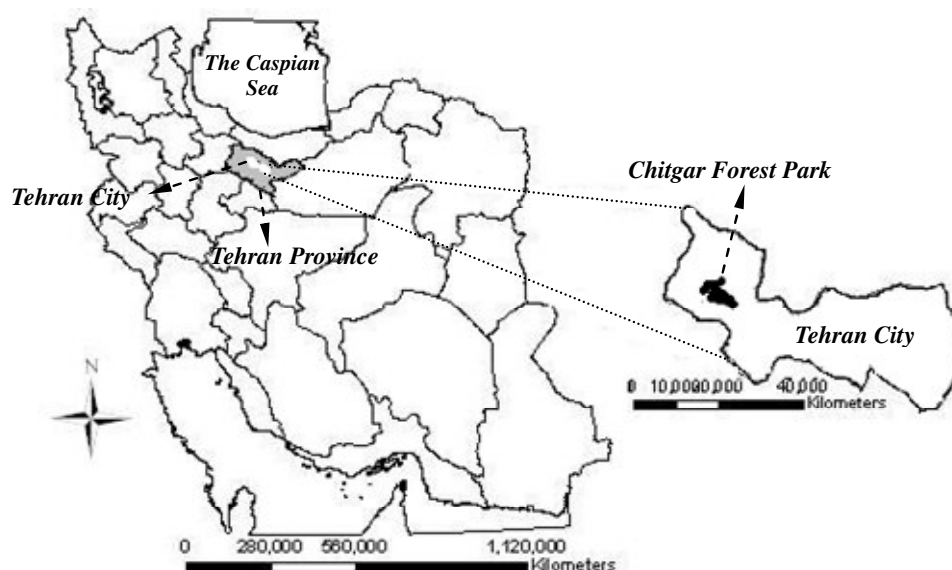


Figure 1. Location of the Chitgar Forest Park near Tehran city, Iran (the sample plot inside the Park not shown).

1305 m asl). For this region, the wettest and driest months are March (45.4 mm; SE: ± 10.7 mm) and August (0.9 mm; SE: ± 0.4 mm), respectively. In this region, the dry period begins in May and ends in October. The wet period extends from November to April, and historically accounts for 88% of the total annual precipitation. The mean annual temperature is 17.2°C (SE: $\pm 0.1^{\circ}\text{C}$); August is the warmest month with average temperature of 29.4°C (SE: $\pm 0.3^{\circ}\text{C}$) and January is the coldest month (3.8°C ; SE: $\pm 0.8^{\circ}\text{C}$).

Field Measurements

Gross Rainfall (P_G)

P_G was measured by 6 cylindrical plastic collectors that were 9 cm in diameter and 20 cm in height. The 6 collectors were placed in a neighboring open area that was 15 m from the *P. eldarica* forest. The quantity of water in the collectors was measured manually using a graduated cylinder with an accuracy of 1 ml. After a rain event, rainwater was measured from each of the rainfall collectors 2 hours following an event if the event occurred during daylight hours or at sunrise if the event occurred at night (Carlyle-Moses *et al.*, 2004).

The average from the 6 rainfall collectors was used to estimate P_G . Rainfall events were defined as separate rain events as long as there was at least 2 hours without rain. In this dry climate, 2 hours was assumed to be sufficient for the canopy to completely dry (Carlyle-Moses *et al.*, 2004).

Throughfall (TF), Stemflow (SF) and Rainfall Interception Loss (I)

TF was measured using 45 rain collectors of the same design as the collectors used to quantify P_G . TF collectors were randomly placed beneath the forest canopy within the study plot (Figure 2). TF volume was measured at the same time P_G was measured. In the present study, SF was not directly measured because *P. eldarica* has rough bark and a canopy structure that is similar to other species with low stemflow (Helvey and Patric, 1965; Geiger, 1965; Llorens *et al.*, 1997, 2000; Lankreijer *et al.*, 1993; Návar, 2012). While *Pinus* trees can have stemflow above 5% of gross precipitation, the vast majority of studies have found *Pinus* species to have stemflow values well below 5% of gross precipitation (Zink, 1967; Steinbuck, 2002). For example, Návar (2011) reported less than 0.20% for

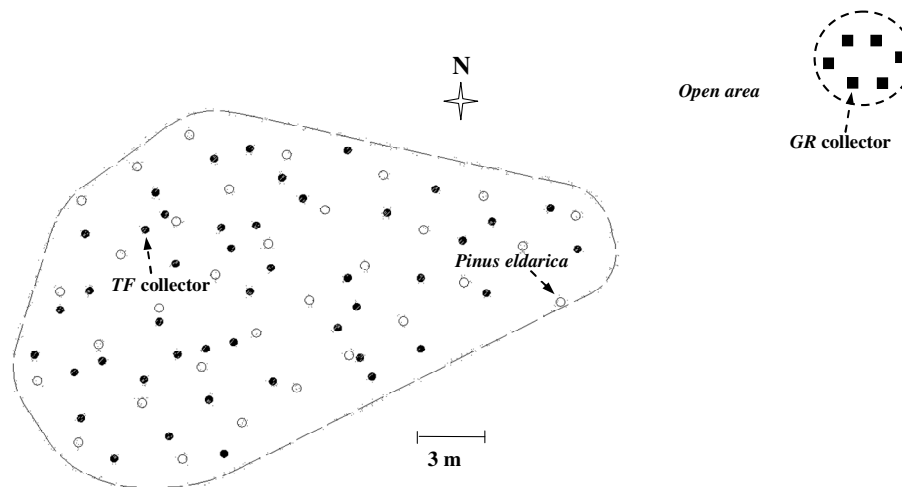


Figure 2. Positions of *Pinus eldarica* trees (open circles) and throughfall (TF) collectors (filled circles) in the study plot. Gross rainfall (P_G) collectors are shown in an open adjacent area. The actual positions of trees and collectors in the study plot were surveyed using a compass and a tape measure.

semi-arid pines of northern Mexico. Furthermore, *P. eldarica* has plagiophile branch structuring. Past research has demonstrated that trees with plagiophile branching have reduced stemflow relative to forests with more erectile branch structure and smooth bark (Pypker et al., 2011, Levia and Frost, 2003, André et al., 2008). Hence, we calculated I as the difference between P_G and TF .

Canopy Storage Capacity (S) and Free Throughfall Coefficient (p)

For the purposes of this paper, we define S as an estimate of the water remaining in the canopy after rainfall ceases and evaporation is negligible (Gash et al., 1995). To estimate S , we applied the commonly used mean method (Jackson, 1975; Pypker et al., 2005; Link et al., 2004; Návar, 2012). The mean method estimates S , p and the ratio of the mean evaporation rate to the mean rainfall intensity (\bar{E}/\bar{R}) by creating two linear regressions (A and B) that relate TF to P_G (Jackson, 1975; Pypker et al., 2005; Link et al. 2004). The first regression line (A) is fit to all the rainfall events where P_G was sufficient to saturate the canopy and the second regression (B) is fit to all the rainfall events where P_G was insufficient to saturate the canopy. When using the mean method, the difference between P_G

and TF at the intersection point provides an estimate of S , the slope of the second regression line provides the estimate of p and one minus the slope of the first regression line represents \bar{E}/\bar{R} (Leyton et al., 1967; Jackson., 1975; Klaassen et al., 1998; Llorens and Gallart, 2000; Pypker et al., 2005). The amount of rainfall sufficient to saturate the canopy (P_s) was estimated subjectively by locating the inflection point on the graph relating TF to P_G .

Gash Model (I)

In this study, a comparison was made between I estimated by the field measurements and by the revised Gash model (Gash et al., 1995). The revised Gash model is a powerful tool for estimating I because of its simple requirements of S , p , and \bar{E}/\bar{R} (Gash et al., 1995). The Gash model is the most common rainfall interception model used in interception studies (Muzylo et al., 2009) although it has been reported by some to incorrectly predict I in sparse forests (Návar, 2012). This model might be a valuable tool for studying rainfall interception loss in forests in Iran. The Gash model is, however, limited by the following assumptions outlined in Gash (1979): (1) rainfall is represented by a series of discrete storms separated by periods long enough to allow the canopy to completely dry up; (2) the

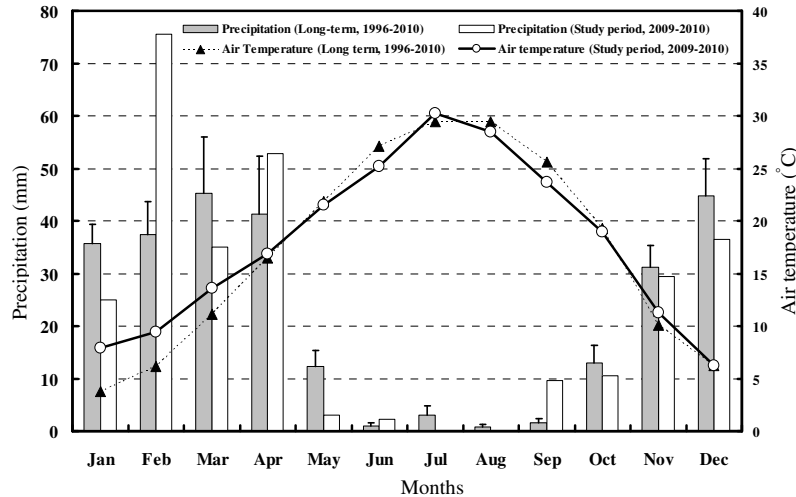


Figure 3. Long-term and study period mean monthly precipitation (mm) and air temperature (°C) at the Chitgar Forest Park Meteorological station. The darker bars and dashed line indicate the long-term (1996-2010) monthly means of the precipitation and air temperature, respectively. The light bars and open circles represent the precipitation and air temperature during the study period, respectively. Error bars show the standard error (SE) of monthly precipitation for the long-term record.

meteorological conditions are constant throughout the storm; and (3) there is no drip from the canopy during wet-up. Clearly, assumptions 2 and 3 are frequently violated during a storm as meteorological conditions such as wind speed, rainfall intensity and vapor pressure deficit can change throughout the storm and wind speeds may vary and shake the canopy causing drip during wet-up. Yet, the Gash model has proven to be very robust in predicting annual I (Gash *et al.*, 1995). The following is the Gash model for sparse canopies (Gash *et al.*, 1995). The interception (I_c) during small storms that were insufficient to saturate the canopy is described by:

$$I_c = c \sum_{j=1}^m P_{G,j} \quad (1)$$

Where, c represents the canopy cover. The canopy cover was assumed to equal 1-p (Gash *et al.*, 1995; Herbst *et al.*, 2008). The amount of interception for n storms sufficient to saturate the canopy (I_n) (i.e. \geq the amount of rainfall to saturate the canopy- P_s) is calculated as sum of the amount of water lost during wet-up (I_w), the evaporation after canopy saturation, but prior to rainfall ceasing (I_s) and the evaporation after the storm ceases (I_a). These interception variables are calculated as:

$$I_w = ncP_s - ncS_c \quad (2)$$

$$I_a = ncS_c \quad (3)$$

$$I_s = (\bar{E}/\bar{R}) \sum_{j=1}^n (P_G - P_s) \quad (4)$$

$$I_n = I_w + I_a + I_s \quad (5)$$

Where, $S_c = S/c$ and (Gash *et al.*, 1995). We parameterized the Gash model using estimates of p , S , P_s , and \bar{E}/\bar{R} by applying the mean method to half of the rainfall data. This resulted in estimates of p , S , P_s , and \bar{E}/\bar{R} equal to 0.16, 2.09 mm, 2.68 mm, and 0.11, respectively. These results were then applied to the other half of the data set to determine if the Gash model can accurately estimate I .

RESULTS

Long-term Average and Observed Meteorology

From September 2009 to April 2010, the cumulative gross precipitation was 279.6 mm, slightly more than the long-term average of 267.6 mm. The annual distribution of precipitation during the study period was similar to that of the long-term average because most of the precipitation recorded



during the study period (91%) occurred during the wet period (from November to May; Figure 3). Historically, the wet period accounts for 88% of the total annual precipitation. The wettest and driest months in the long-term records were March (45.4 mm) and August (0.9 mm), respectively. However, during the study period, the wettest month was February (75.5 mm) and the driest months were July and August (0 mm) (Figure 3). Average air temperature was 17.8°C during the study period, similar to the long-term average temperature of 17.2°C. Historically, August was the warmest month (average temperature of 29.4°C), but, during the study period, July was the warmest (30.3°C). Like the long-term average, January was the coldest month during the study period (7.9°C) (Figure 3).

Gross Rainfall (P_G) and Throughfall (TF)

From September 2009 to April 2010, 164.8 mm of rain fell in thirty rainfall events. Cumulative P_G for an individual rainfall event ranged from 0.2 mm to 27.5 mm, with a mean P_G depth per event of 5.5 mm. The rainfall events were grouped into three classes (Table 1) to allow for a better understanding of the relationship between P_G and I . The three classes are: $P_G \leq 1.5$ mm, $1.5 \text{ mm} < P_G \leq 5.5$ mm, and $P_G > 5.5$ mm. This resulted in 10 rainfall events being allocated to each of the three classes.

Of the 30 rainfall events recorded during the measurement period, 103.6 mm, or 62.8% of the cumulative P_G reached the forest floor as TF . Mean TF was 3.5 mm (CV: 72.9%) or 38.6% of P_G and ranged from 0.0 to 80.5% for P_G events of 0.2 to 27.5 mm, respectively. For the $P_G \leq 1.5$ mm, $1.5 \text{ mm} \leq P_G \leq 5.5$ mm, and $P_G \geq 5.5$ mm rainfall event classes, the mean $TF:P_G$ values were 8.9, 40, and 67%, respectively (Table 1).

Rainfall Interception Loss (I)

Rainfall interception loss totaled 61.2 mm, or 37.2% of the total P_G for the study period. The percentage lost to I depended on storm size, with the percentage varying from 19.5% of P_G for larger rainfall events (27.5 mm) to 100.0% of P_G (0.2 mm) for smaller rainfall events.

The ratio of I to P_G (relative I or $I:P_G$) was correlated with P_G (Figure 4). The mean values of $I:P_G$ showed a decreasing trend, with a decreasing ratio as P_G increased. A negative logarithmic significant relationship ($r^2 = 0.861$; $P \text{ value} \leq 0.01$) was fitted between $I:P_G$ and P_G .

Canopy (Water) Storage Capacity (S)

S was determined using a method that related P_G to TF (Figure 5). The mean method (Klaassen *et al.*, 1998; Jackson, 1975) estimated S to be 1.8 mm (Figure 5).

Lastly, the mean method estimated p and \bar{E}/\bar{R} to be 0.1 and 0.16, respectively.

Gash Model

The Gash Model accurately estimated I (Figure 6). On a per-storm basis, the difference between modeled and measured values ranged from 0.03 to 1.6 mm, with the greatest differences occurring for the largest storms. The Gash model predicted I to be 26.7 mm, and this only differed from the measured I (25.6 mm) by 1.1 mm. The slope of the line relating the measured and modeled estimates of I was 0.86 (95% CI = 0.17). The measured and modeled estimates of I were not statistically different because the slope was not significantly different from 1 ($P\text{-value} > 0.1$).

Table 1. Mean cumulative gross rainfall (P_G), mean throughfall as a percentage of P_G ($TF:P_G$), and mean rainfall interception as a percentage of P_G ($I:P_G$) for rainfall events from September 2009 to April 2010. The relationships are provided for three P_G classes.

P_G class (mm)	Frequency	P_G (mm)	$TF:P_G$ (%)	$I:P_G$ (%)
< 1.5	10	7.8	8.9	91.1
1.5-5.5	10	34.7	40.0	60.0
> 5.5	10	122.3	67.0	33.0

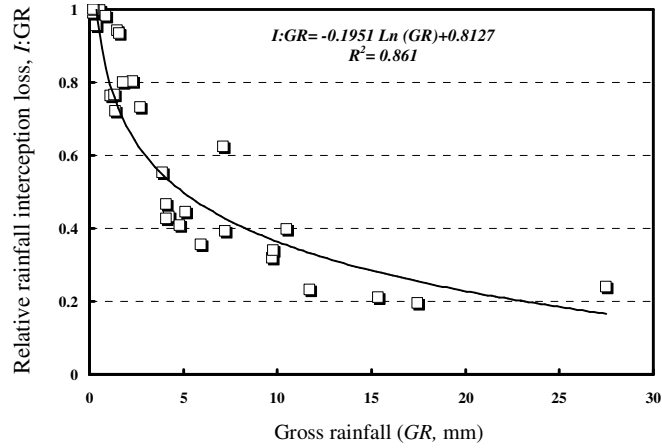


Figure 4. The relationship between relative rainfall interception loss ($I:P_G$) and P_G in *Pinus eldarica* afforestation.

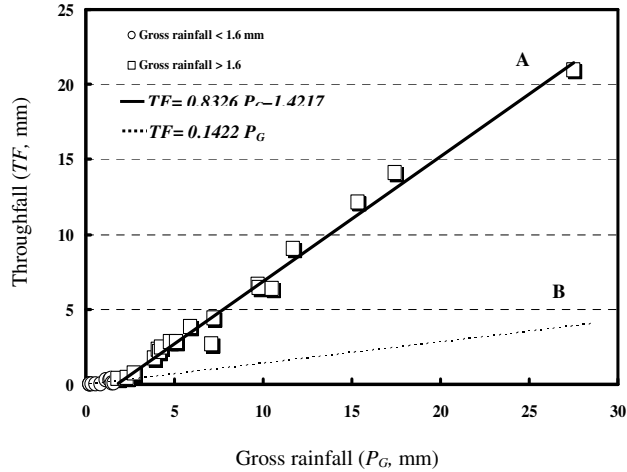


Figure 5. Estimation of canopy storage capacity (S) and free throughfall coefficient (p) using the relationship between gross rainfall (P_G) and throughfall (TF) during the study period (Mean method). The difference between P_G and TF at the inflection point estimates the canopy storage capacity (S). The slope of regression B represents p and 1 minus the slope of the A regression line represents the ratio of mean evaporation to mean rainfall intensity (\bar{E} / \bar{R}).

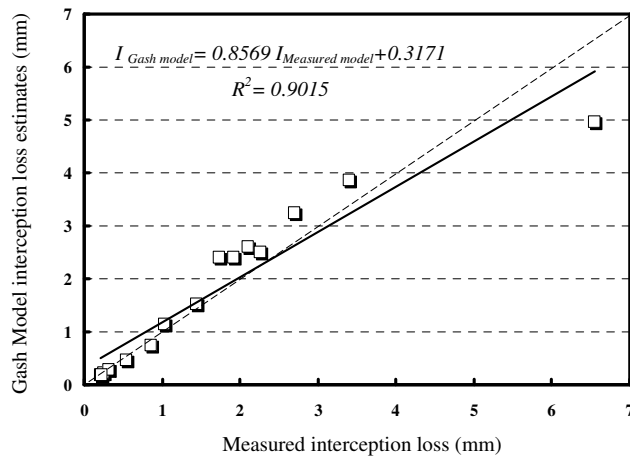


Figure 6. A comparison between Gash model estimates of I to measured I for half of the data. The 1:1 line is the dashed line. The solid line represents a linear regression between the x-axis and y-axis.



DISCUSSION

Partitioning of Gross Precipitation between Throughfall and Interception Loss

A review of the literature on rainfall partitioning in various pine stands (Table 2) indicates that the values for $TF:P_G$ and $I:P_G$ obtained in the present study differed slightly with those measured in other pine forests (Table 3). Llorens *et al.* (1997) reported that the average values of $TF:P_G$ and $I:P_G$ in a *Pinus sylvestris* forest in Eastern Pyrenees, Spain, were 0.747 and 0.24, respectively. Mahendrapa (1990) reported TF and I for a *Pinus strobus* plantation in Canada to be 65 and 30.7% of annual P_G , respectively. In Portugal, the measured value of TF in a *Pinus pinaster* forest was 82.6% of P_G during the two years of measurement (Valente *et al.*, 1997). It is noteworthy that value of $I:P_G$ obtained in our study (0.37) was on the high end of the 0.12 to 0.42 measured by others in needle-leaved evergreen forests (Hibbert, 1967; Zinke, 1967) (Table 2). The partitioning of rainfall into TF and I in forest ecosystems has been demonstrated to be a function of incident rainfall characteristics (amount, intensity, duration, and temporal distribution of rainfall events) (Jackson, 1975; Crockford and Richardson, 2000; Xiao *et al.*, 2000; Marin *et al.*, 2000; Huber and Iroumé, 2001; Iroumé and Huber, 2002; Link *et al.*, 2004; Návar, 2012), meteorological conditions (air temperature, relative humidity, wind speed, and wind direction) (Crockford and Richardson, 2000), and forest structure (species composition, stand age, basal area, stand density and canopy morphology and architecture) (Forgeard *et al.*, 1980; Xiao *et al.*, 2000; Iroumé and Huber, 2002; Carlyle-Moses, 2004; Fleischbein *et al.*, 2005; Deguchi *et al.*, 2006; Staelens *et al.*, 2008; Muzyllo *et al.*, 2009). Given the dry nature of this region, it is likely that the differences in rainfall partitioning reported by other researchers were due, in part, to differences in the above mentioned factors.

The size of P_G had a major impact on the partitioning of rainfall into TF and I for the *P. eldarica* afforestation in this study. As the size

of P_G increased, the ratio of I to P_G ($I:P_G$) decreased. For example, as P_G increased from < 1.5 mm to > 5.5 mm, $I:P_G$ decreased from 0.91 to 0.33, respectively (Table 1). In this study, 67% of the rainfall events were less than 5.5 mm, therefore, the large proportion of incident P_G wetted the crown surface and subsequently contributed to interception loss. Therefore, part of the difference in interception loss between this study and others in the literature may have resulted from different storm sizes. However, while TF and P_G were well correlated, the lowest TF values were not synchronized with the lowest values of P_G (0.22), thereby suggesting that the climatic factors also played a very important role in the rainfall partitioning.

Estimate of Canopy (Water) Storage Capacity

In the present study, S was estimated to be 1.8 mm (Table 3). The estimates of S in the present study fall within the range reported for other pine forests. For example, using artificial wetting, Llorens and Gallart (2000) determined S in a *Pinus sylvestris* forest (1,400 trees ha⁻¹) in Eastern Pyrenees, Spain, to be 2 mm. Liu (1998) determined S to be 0.7 mm for a *Pinus elliottii* afforestation (1,190 trees ha⁻¹ similar to this study). Llorens (1997) also reports an S of 1.3 mm by an indirect method for a *Pinus sylvestris* afforestation (2,400 trees ha⁻¹). Lastly, Návar (2012) reported a mean value of 1.03 mm for temperate, dry *P. pseudostrobus* forests in northern Mexico.

Changes in rainfall intensity (Calder *et al.*, 1996; Jackson, 1975) and wind speed (Jackson, 1975; Hörmann *et al.*, 1996) are correlated with changes in S . The nature of the intercepting surface also controls the size of S (type of species, leaf shape, dimension and orientation) (Jackson, 1975; Liu, 1998; Llorens and Gallart, 2000; Fleischbein *et al.*, 2005). Wind movement of the canopy may greatly reduce the amount of water which can be held before drainage occurs (Jackson, 1975; Hörmann *et al.*, 1996). Therefore, the amount of S may vary from event to event even if canopy characteristics remain constant.

Table 2. A review of measured values of relative throughfall ($TF:P_G$), relative rainfall interception ($I:P_G$) as well as relative stemflow ($SF:P_G$) obtained from various research carried out on different species of pine (stand). TF , SF , I , and P_G are referred to throughfall, stemflow, rainfall interception, and gross rainfall, respectively.

Tree Species	$I:P_G$ (%)	$TF:P_G$ (%)	$SF:P_G$ (%)	Tree density (Stem ha ⁻¹)	Location	References
<i>Pinus pinaster</i>	17.1	82.6	0.3	312	Portugal	Valente <i>et al.</i> (1997)
<i>Pinus radiata</i>	18.3	72.8	8.9	1708	South-Eastern Australia	Crockford and Richardson (1990)
<i>Pinus wallichiana</i>	21	76.3	2.7	1200	India	Singh (1987)
<i>Pinus densiflora</i>	14	83	3	1575	Etajima Island (West Japan)	Mitsudera <i>et al.</i> (1984)
<i>Pinus radiata</i>	30	----	----	450	New Zealand	Kelliher <i>et al.</i> (1992)
<i>Pinus sylvestris</i>	32	38-53	15-30	4600	Crowthorne, Berks	Rutter (1963)
<i>Pinus massoniana</i>	27.2	70.4	2.4	2628	China	Cao <i>et al.</i> (2008)
<i>Pinus sylvestris</i>	24	74.7	1.3	2400	Eastern Pyrenees, Spain	Llorens <i>et al.</i> (1997)
<i>Pinus radiata</i>	26.5	----	----	1493	Shoalhaven-Australia	Pook <i>et al.</i> (1991)
<i>Pinus nigra</i>	35	65	----	600	Southeast England	Rutter <i>et al.</i> (1971)
<i>Pinus sylvestris</i>	42.4	57.6	----	1870	Northeast Scotland	Gash <i>et al.</i> (1980)
<i>Pinus strobus</i>	30.7	65	5.3	----	Canada	Mahendrappa (1990)
<i>Pinus resinosa</i>	28.3	69	0.7	----	Canada	Mahendrappa (1990)
<i>Pinus pinaster</i>	12.6 - 21	76-83	1-6	800	Southwest Europe	Loustau <i>et al.</i> (1992)
<i>Pinus pinaster</i>	12.5	87.5	----	430	France	Lankreijer <i>et al.</i> (1993)
<i>Pinus pseudostrabus</i>	17.8	82.0	0.20	125	Northern Mexico	Návar (2011; 2012)

Table 3. A review of measured values of canopy storage capacity (S), free throughfall coefficient (p) and the measurement methods used to estimate S and p in coniferous forests.

Species	Density (trees ha ⁻¹)	S (mm)	p	Method	References
<i>Pinus sylvestris</i>	509	2.3	0.2	Direct	Llorens and Gallart (2000)
<i>Pinus sylvestris</i>	764	1.2	0.4	Direct	Llorens and Gallart (2000)
<i>Pinus sylvestris</i>	800	0.8	0.3	Regression line	Gash and Morton (1978)
<i>Pinus sylvestris</i>	1400	2	0.2	Direct	Llorens and Gallart (2000)
<i>Pinus sylvestris</i>	1782	1.5	0.3	Direct	Llorens and Gallart (2000)
<i>Pinus sylvestris</i>	1870	1	0.1	Regression line	Gash <i>et al.</i> (1980)
<i>Pinus sylvestris</i>	2400	1.3	0.1	Regression line	Llorens (1997)
<i>Pinus sylvestris</i>	2674	2.7	0.1	Direct	Llorens and Gallart (2000)
<i>Pinus sylvestris</i>	2900	0.3	----	----	Perttu <i>et al.</i> (1980)
<i>Pinus sylvestris</i>	4600	1.6	----	Regression line	Rutter (1963)
<i>Pinus eldarica</i>	1185	1.77	0.1	Regression line (Mean method)	This study
<i>Pinus pinaster</i>	312	0.4	0.4	Regression line	Valente <i>et al.</i> (1997)
<i>Pinus pinaster</i>	430	0.3	0.4	Regression line	Lankreijer <i>et al.</i> (1993)
<i>Pinus pinaster</i>	800	0.5	0.6	Regression line	Loustau <i>et al.</i> (1992)
<i>Pinus elliottii</i>	464	0.4	---	Direct	Liu (1998)
<i>Pinus elliottii</i>	496	0.5	---	Direct	Liu (1998)
<i>Pinus elliottii</i>	672	0.4	---	Direct	Liu (1998)
<i>Pinus elliottii</i>	1190	0.7	---	Direct	Liu (1998)
<i>Pinus nigra</i>	600	1.1	0.3	Regression line	Rutter <i>et al.</i> (1971)
<i>Pinus nigra</i>	600	1	----	----	Robins (1974)
<i>Pinus radiata</i>	450	0.4	----	Regression line	Kelliher <i>et al.</i> (1992)
<i>Pinus pseudostrabus</i>	125	1.03	0.26	Regression line and optical densitometers	Návar (2012)



Several studies have investigated the effects of changes in forest cover on the water recharge (Bosch and Hewlett, 1982; Blackie, 1993; Sahin and Hall, 1996). The reduction in water recharge upon afforestation is mainly due to an increase in interception, evaporation, and transpiration (Van der Salm *et al.*, 2007). The rainfall interception from afforestations with *P. eldarica* in Iran is considerable, averaging 37% of P_G . Therefore, rainfall interception loss needs to be considered in future water balance studies and in the selection of tree species for afforestation practices. Furthermore, future research is needed to quantify the full hydrological (transpiration and rainfall interception loss) effect of these afforestation practices.

Gash Model

The Gash model proved to accurately estimate I for storms up to 30 mm in size. Past research on I in other coniferous forests have also successfully applied the Gash model (Gash *et al.*, 1980; Gash *et al.*, 1999; Valente *et al.*, 1997; Návar *et al.*, 1999a and b; Návar, 2013). As with past research, there were some inaccuracies on an individual storm basis (e.g. Llorens, 1997). However, the model often accurately predicts total I (e.g. Llorens, 1997). Hence, his model could be a valuable tool for estimating TF in watersheds dominated by *P. eldarica* for years with varying annual precipitation. Ultimately, the estimates of TF could be used to better understand annual changes in streamflow.

CONCLUSIONS

This study was carried out in a forty-year-old *Pinus eldarica* afforestation during a 9-month period in a semi-arid climate zone of Iran. I and TF represented 37.2 and 62.8% of annual P_G , respectively. It was observed that rainfall partitioning into TF and I was strongly affected by the size of P_G ; with the ratio of $I:P_G$ declining as P_G increased. S in the *P. eldarica* stand was 1.8 mm.

This research is the first to document rainfall partitioning and S in a *P. eldarica*

afforestation. In the semi-arid climate zone of Iran, plant growth and productivity is strongly affected by water availability. Therefore, I should be considered when selecting species for afforestation projects in the semi-arid climate regions as it can be significant.

REFERENCES

1. André, F., Mathieu, J. and Ponette, Q. 2008. Effects of Biological and Meteorological Factors on Stemflow Chemistry within a Temperate Mixed Oak-beech Stand. *Sci. Total. Environ.*, **393**: 72-83.
2. Asdak, C., Jarvis, P. G., van Gardingen, P. and Frazer, A. 1998. Rainfall Interception Loss in Unlogged and Logged Forest Areas of Central Kalimantan, Indonesia. *J. Hydrol.*, **206**: 237-244.
3. Blackie, J. R. 1993. The Water Balance of the Balquhiddier Catchments. *J. Hydrol.*, **145**: 239-257.
4. Bosch, J. M. and Hewlett, J. D. 1982. A Review of Catchment Experiments to Determine the Effect of Vegetation Changes on Water Yield and Evapotranspiration. *J. Hydrol.*, **55**: 323-3-23.
5. Calder, I. R., Hall, R. L., Rosier, P. T. W., Bastable, H. G. and Prasanna, K. T. 1996. Dependence of Rainfall Interception on Drop Size. 2. Experimental Determination of the Wetting Functions and Two-layer Stochastic Model Parameters for Five Tropical Tree Species. *J. Hydrol.*, **185**: 379-388.
6. Cao, Y., Ouyang, Z. Y., Zheng, H., Huang, Z. G., Wang, X. K. and Miao, H. 2008. Effects of Forest Plantations on Rainfall Redistribution and Erosion in the Red Soil Region of Southern China. *Land Degrad. Develop.* **19**: 321-330.
7. Carlyle-Moses, D. E., 2004. Throughfall, Stemflow, and Canopy Interception Loss Fluxes in a Semi-arid Sierra Madre Oriental Matorral Community. *J. Arid. Environ.*, **58**: 181-202.
8. Carlyle-Moses, D. E., Flores Laureano, J. S. and Price, A. G. 2004. Throughfall and Throughfall Spatial Variability in Madrean Oak Forest Communities of Northeastern Mexico. *J. Hydrol.*, **297**: 124-135.
9. Crockford, R. H. and Richardson, D. P. 1990. Partitioning of Rainfall in a Eucalypt Forest and Pine Plantation in Southeastern Australia. IV. The Relationship of Interception and Canopy Storage Capacity, the Interception of These Forests, and the Effect on Interception of Thinning the Pine Plantation. *Hydrol. Process.*, **4**: 169-188.
10. Crockford, R. H. and Richardson, D. P. 2000. Partitioning of Rainfall into Throughfall,

- Stemflow and Interception: Effect of Forest Type, Ground Cover and Climate. *Hydrol. Process.*, **14**: 2903–2920.
11. De los Rios-Carrasco, E., De Hoogh, R. and Návar, J. 2009. Projections of Carbon Stocks in Sites Reforested with Pinyon Pine Species in Northeastern Mexico. *Arid Land Res. Manag.*, **23**: 342–358.
 12. Deguchi, A., Hattori, S. and Park, H. 2006. The Influence of Seasonal Changes in Canopy Structure on Interception Loss: Application of the Revised Gash Model. *J. Hydrol.*, **319**: 80–102.
 13. Fleischbein, K., Wilcke, W., Goller, R., Boy, J., Valarezo, C., Zech, W. and Knoblich, K. 2005. Rainfall Interception in a Lower Montane Forest in Ecuador: Effects of Canopy Properties. *Hydrol. Process.*, **19**(7): 1355–1371.
 14. Forgeard, F., Gloaguen, J. C. and Touffet, J. 1980. Interception des Précipitations et Apports au sol D'éléments Minéraux par les eaux de Pluie et les Pluviolessivats dans une Hêtraie Atlantique et dans Quelques Peuplements Résineux de Bretagne. *Ann. For. Sci.*, **37**: 53–71.
 15. Gash, J. H. C. and Morton, A. J. 1978. An Application of the Rutter Model to the Estimation of the Interception Loss from Thetford Forest. *J. Hydrol.*, **38**: 49–58.
 16. Gash, J. H. C. 1979. An Analytical Model of Rainfall Interception by Forest. *Quart. J. Roy. Meteorol. Soc.*, **105**: 43–55.
 17. Gash, J. H. C., Wright, I. R. and Lloyd, C. R. 1980. Comparative Estimates of Interception Loss from Three Coniferous Forests in Great Britain. *J. Hydrol.*, **48**: 89–105.
 18. Gash, J.H.C., Lloyd, C.R., and Lachaud, G. 1995. Estimating Sparse Forest Rainfall Interception with an Analytical Model. *J. Hydrol.*, **170**: 79–86.
 19. Gash, J. H. C., Valente, F. and David, J. S. 1999. Estimates and Measurements of Evaporation from Wet, Sparse Pine Forest in Portugal. *Agric. For. Meteorol.*, **94**: 149–158.
 20. Geiger, R. 1965. *The Climate near the Ground*. Harvard University Press, Cambridge, Massachusetts, 611PP.
 21. Grünzweig, J. M., Lin, T., Rotenberg, E., Schwartz, A. and Yakir, D. 2003. Carbon Sequestration in Arid-land Forest. *Global Change Biol.*, **9**: 791–799.
 22. Helvey, J. D. and Patric, J. H. 1965. Design Criteria for Interception Studies. In: "Design of Hydrological Networks; Proceedings of a Symposium", June 1965, Quebec City Canada. *Int. Assoc. Sci. Hydrol.*, **67**: 131–137.
 23. Herbst, M., Rosier, P. T. W., McNeil, D. D., Harding, R. J. and Gowing, D. J. 2008. Seasonal Variability of Interception Evaporation from the Canopy of a Mixed Deciduous Forest. *Agric. For. Meteorol.*, **148**: 1655–1667.
 24. Hibbert, A. R. 1967. Forest Treatment Effects on Water Yield. In: "Proceedings International Symposium on Forest Hydrology", (Eds.): Sopper, W. E. and Lull, H. W.. Pergamon Press, Oxford, PP. 527–544.
 25. Hörmann, G., Branding, A., Clemen, T., Herbst, M., Hinrichs, A. and Thamm, F. 1996. Calculation and Simulation of Wind Controlled Canopy Interception of a Beech Forest in Northern Germany. *Agric. For. Meteorol.*, **79**: 131–148.
 26. Huber, A. and Iroumé, A. 2001. Variability of Annual Rainfall Partitioning for Different Sites and Forest Cover in Chile. *J. Hydrol.*, **248**: 78–92.
 27. Hutchinson, I. and Roberts, M. C. 1981. Vertical Variation in Stemflow Generation. *J. Appl. Ecol.*, **18**: 521–527.
 28. Hutjes, R., Wierda, A. and Veen, A. 1990. Rainfall Interception in the Tai Forest, Ivory Coast: Application of Two Simulation Models to a Humid Tropical System. *J. Hydrol.*, **114**: 259–275.
 29. Hüttl, R. F., Schneider, B. U. and Farrell, E. P. 2000. Forests of the Temperate Region: Gaps in Knowledge and Research Needs. *For. Ecol. Manag.*, **132**: 83–96.
 30. Iroumé, A. and Huber, A. 2002. Comparison of Interception Losses in a Broadleaved Native Forest and a *Pseudotsuga menziesii* (Douglas Fir) Plantation in the Andes Mountains of Southern Chile. *Hydrol. Process.*, **16**: 2347–2361
 31. Jackson, I. J. 1975. Relationships between Rainfall Parameters and Interception by Tropical Rainforest. *J. Hydrol.*, **24**: 215–238.
 32. Kelliher, F. M., Whitehead, D. and Pollock, D. S. 1992. Rainfall Interception by Trees and Slash in a Young *Pinus radiata* D. Don Stand. *J. Hydrol.*, **131**: 187–204.
 33. Klaassen, W., Bosveld, F. and de Water, E. 1998. Water Storage and Evaporation as Constituents of Rainfall Interception. *J. Hydrol.*, **212–213**: 36–50.
 34. Lankreijer, H. J. M., Hendriks, M. J. and Klaassen, W. 1993. A Comparison of Models Simulating Rainfall Interception of Forests. *Agric. For. Meteorol.*, **64**: 187–199
 35. Levia, D. and Frost, E. 2003. Variability of Throughfall Volume and Solute Inputs in Wooded Ecosystems. *Prog. Phys. Geogr.*, **30**: 605–632
 36. Leyton, L., Reynolds, E. R. C. and Thompson, F. B. 1967. Rainfall Interception in Forest and Moorland. In: "International Symposium on Forest Hydrology", (Eds.): Sopper, W. E. and



- Lull, H. W.. Pergamon Press, Pennsylvania State University, PP. 163–178.
37. Link, T. E., Unsworth, M. and Marks, D. 2004. The Dynamics of Rainfall Interception by a Seasonal Temperate Rainforest. *Agric. For. Meteorol.*, **124**: 171–191.
38. Liu, S. G. 1997. A New Model for the Prediction of Rainfall Interception in Forest Canopies. *Ecol. Model.*, **99**: 151–159.
39. Liu, S.G., 1998. Estimation of Rainfall Storage Capacity in the Canopies of Cypress Wetlands and Slash Pine Uplands in North-Central Florida. *J. Hydrol.*, **207**: 32–41.
40. Llorens, P. 1997. Rainfall Interception by a *Pinus sylvestris* Forest Patch Overgrown in a Mediterranean Mountainous Abandoned Area. II. Assessment of the Applicability of Gash's Analytical Model. *J. Hydrol.*, **199(3-4)**: 346–359.
41. Llorens, P., Poch, R., Latron, J. and Gallart, F. 1997. Rainfall Interception by *Pinus Sylvestris* Forest Patch Overgrown in a Mediterranean Mountainous Abandoned Area. I. Monitoring Design and Results to Event Scale. *J. Hydrol.*, **199(3-4)**: 331–345.
42. Llorens, P. and Gallart, F. 2000. A Simplified Method for Forest Water Storage Capacity Measurement. *J. Hydrol.*, **240**: 131–144.
43. Loustau, D., Bergigier, P. and Granier, A. 1992. Interception Loss, Throughfall and Stemflow in a Maritime Pine Stand. II. An Application of Gash's Analytical Model of Interception. *J. Hydrol.*, **138**: 469–485.
44. Lloyd, C., Gash, J., Shuttleworth, W. and Marques, FA. 1988. The Measurement and Modelling of Rainfall Interception by Amazonian Rain Forest. *Agric. For. Meteorol.*, **43**: 277–294.
45. Mahendrapa, M. K. 1990. Partitioning of Rainwater and Chemical into Throughfall and Stemflow in Different Forest Stands. *For. Ecol. Manag.*, **30**: 65–72.
46. Marin, C. T., Bouten, W. and Sevink, J. 2000. Gross Rainfall and Its Partitioning into Throughfall, Stemflow and Evaporation of Intercepted Water in Four Forest Ecosystems in Western Amazonia. *J. Hydrol.*, **237**: 40–57.
47. Mitsudera, M., Kamata, Y. and Nakane, K. 1984. Effect of Fire on Water and Major Nutrient Budgets in Forest Ecosystems III. Rainfall Interception by Forest Canopy. *Jap. J. Ecol.*, **34**: 15–25.
48. Muzylo, A., Llorens, P., Valente, F., Keizer, J. J., Domingo, F. and Gash, J. H. C. 2009. Review of Rainfall Interception Modelling. *J. Hydrol.*, **370**: 191–206.
49. Nívar, J. and Bryan, R. 1990. Interception Loss and Rainfall Redistribution by Three Semiarid Growing Shrubs in Northeastern Mexico. *J. Hydrol.*, **115**: 51–63.
50. Nívar, J., Charles, F. and Jurado, E. 1999a. Spatial Variations of Interception Loss Components by *Tamaulipan thornscrub* in Northeastern Mexico. *For. Ecol. Manage.*, **124**: 231–239.
51. Nívar, J., Carlyle-Moses, E. and Martínez, A. 1999b. Interception Loss from the *Tamaulipan thornscrub* of Northeastern Mexico: An Application of the Gash Analytical Interception Loss Model. *J. Arid Environ.*, **41**: 1–10.
52. Nívar, J. 2011. Stemflow variation in Mexico's Northeastern Forest Communities: Its Contribution to Soil Moisture Content and Aquifer Recharge. *J. Hydrol.*, **408**: 35–42.
53. Nívar J. 2013. The Performance of the Reformulated Gash's Interception Loss Model in Mexico's Northeastern Temperate Forests. *Hydrol. Proces.*, 1626–1633.
54. Pearce, A., Gash, J. and Stewart, J. 1980. Rainfall Interception in a Forest Stand Estimated from Grassland Meteorological Data. *J. Hydrol.*, **46**: 147–163.
55. Perttu, K., Bishop, W., Grip, H., Jansson, P. E., Lindgren, A., Lindroth, A. and Noren, B. 1980. Micrometeorology and Hydrology of Pine Forest Ecosystems. I. Field Studies. Structure and Function of Northern Coniferous Forest: An Ecosystem Study. *Ecol. Bull.*, **32**: 75–121.
56. Pook, E. W., Moore, P. H. R. and Hall, T. 1991. Rainfall Interception by Trees of *Pinus radiata* and *Eucalyptus viminalis* in a 1300 mm Rainfall Area of Southeastern New South Wales. I. Gross Losses and Their Variability. *Hydrol. Processes.*, **5**: 127–141.
57. Pypker, T. G., Bond, B. J., Link, T. E., Marks, D. and Unsworth, M. H. 2005. The Importance of Canopy Structure in Controlling the Interception Loss of Rainfall: Examples from a Young and an Old-grown Douglas-fir Forest. *Agric. For. Meteorol.*, **130**: 113–129.
58. Pypker, T. G., Levia, D. F., Staelens, J. and Van Stan II, J. T. 2011. Chapter XVII Canopy Structure in Relation to Hydrological and Biogeochemical Fluxes. In: "*Forest Hydrology and Biogeochemistry: Synthesis of Past Research and Future Directions*", (Eds.): Levia, D. F., Carlyle-Moses, D. E. and Tanaka, T. Springer-Verlag, Heidelberg, Germany, 371–378.
59. Robins, P. C. 1974. A Method of Measuring the Aerodynamic Resistance to the Transport of Water Vapour from Forest Canopies. *J. Appl. Ecol.*, **11**: 315–325.
60. Rutter, A. J. 1963. Studies in the Water Relations of *Pinus sylvestris* in Plantation Conditions. I. Measurements of Rainfall and Interception. *J. Ecol.*, **51**: 315–325.

61. Rutter, A. J., Kershaw, K. A., Robins, P. C. and Monton, A. J. 1971. A Predictive Model of Rainfall Interception Forests. 1. Derivation of the Model from Observations in a Plantation of Corsican Pine. *Agric. Meteorol.* **9**: 367-384.
62. Sahin, V. and Hall, M. J. 1996. The Effects of Afforestation and Deforestation on Water Yields. *J. Hydrol.*, **178**: 293-309.
63. Sardabi, H. 1998. *Eucalypt and Pine Species Trials on the Caspian Littoral and Lowlands of Iran*. Technical Publication No. 193, Research Institute of Forests and Rangelands (RIFRI), Tehran, Iran.
64. Scatena, F. N. 1990. Watershed Scale Rainfall Interception on Two Forested Watersheds in the Luquillo Mountains of Puerto Rico. *J. Hydrol.*, **113**: 89-102.
65. Singh, R. P. 1987. Rainfall Interception by *Pinus Wallichiana* Plantation in Temperate Region of Himachal Pradesh, India. *Indian For.* 559-566.
66. Staelens, J., Schrijver, A. D., Verheyen, K. and Verhoest, N. 2008. Rainfall Partitioning into Throughfall, Stemflow, and Interception within a Single Beech (*Fagus sylvestris* L.) Canopy: Influence of Foliation, Rain Event Characteristics, and Meteorology. *Hydrol. Process.*, **22**: 33-45.
67. Steinbuck, E. 2002. The Influence of Tree Morphology on Stemflow in a Redwood Region Second- Growth Forest. MSc. Thesis, California State University, Chico, CA, 55 PP.
68. Teklehaimanot, Z. and Jarvis, P. G. 1991. Direct Measurement of Evaporation of Intercepted Water from Forest Canopies. *J. Appl. Ecol.*, **28**: 603-618.
69. Tobón Marin, C., Bouten, W. and Sevink, J. 2000. Gross Rainfall and Its Partitioning into Throughfall, Stemflow and Evaporation of Intercepted Water in Four Forest Ecosystems in Western Amazonia. *J. Hydrol.*, **237**: 40-57.
70. Valente, F., David, J. S. and Gash, J. H. C. 1997. Modelling Interception Loss for Two Sparse Eucalypt and Pine Forests in Central Portugal Using Reformulated Rutter and Gash Analytical Models. *J. Hydrol.*, **190**: 141-162.
71. Van der Salm, C., Rosenqvist, L., Vesterdal, L., Hansen, K., Denier van der gon, H., Bleeker, A., Wieggers, R. and Van der torn, A. 2007. Interception and Water Recharge Following Afforestation: Experiences from Oak and Norway Spruce Chronosequences in Denmark, Sweden and the Netherlands. *Environmental Effects of Afforestation in North-Western Europe*: 53-77.
72. Van Dijk, A. and Bruijnzeel, L. 2001. Modelling Rainfall Interception by Vegetation of Variable Density Using an Adapted Analytical Model. Part 1. Model Description. *J. Hydrol.*, **247**: 230-238.
73. Xiao, Q., McPherson, E. G., Ustin, S. L., Grismer, M. E. and Simpson, J. R. 2000. Winter Rainfall Interception by Two Mature Open-grown Trees in Davis, California. *Hydrol. Process.*, **14**: 763-784.
74. Zhou, G. Y., Wei, X. H. and Yan, J. H. 2002. Impacts of Eucalyptus (*Eucalyptus exserta*) Plantation on Sediment Yield in Guangdong Province, Southern China. A Kinetic Energy Approach. *Catena*, **49**: 231-251.
75. Zinke, P. J. 1967. Forest Interception Study in the United States. In: "Forest Hydrology" (Eds.): Sopper, W. E. and Lull, H. W.. Pergamon, Oxford, PP. 137-161.

اندازه گیری باران ربایی کاج تهران (*Pinus eldarica*) در اقلیم نیمه خشک: کاربرد

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چکیده

باران ربایی (I)، ظرفیت نگهداری آب روی تاج پوشش (S) و نسبت میانگین تبخیر به میانگین شدت بارندگی (\bar{E}/\bar{R}) از اجزای مهم تعادل آب در مناطق خشک و نیمه خشک می باشند. هدف از این تحقیق تعیین مقدار باران ربایی و ظرفیت نگهداری آب روی تاج پوشش و همچنین ارزیابی مدل Gash



برای بر آورد باران‌ریایی در یک توده خالص جنگلکاری شده کاج تهران (*Pinus eldarica*) (Medw) واقع در پارک جنگلی چیتگر تهران می‌باشد. اندازه‌گیری بارندگی در هر بارش (PG) و تاج-بارش (TF) از مهر ۱۳۸۸ تا اردیبهشت ۱۳۸۹ صورت گرفت. در این دوره در مجموع ۱۶۴/۸ میلی‌متر بارندگی جمع‌آوری شد که ۶۱/۲ میلی‌متر آن به باران‌ریایی اختصاص پیدا کرد. باران‌ریایی از تفاوت بین بارندگی در هر بارش و تاج‌بارش محاسبه گردید. نسبت باران‌ریایی به بارندگی در هر بارش ($I:PG$) بین ۰/۱۹۵ و ۱ و به طور میانگین ۰/۶۱۴ بدست آمد. نتایج نشان داد رابطه لگاریتمی قوی بین نسبت باران‌ریایی به بارندگی و بارندگی در هر بارش ($r^2 = 0.861$; $p \text{ value} \leq 0.01$) وجود دارد به طوری که با افزایش مقدار بارندگی در هر بارش، نسبت باران‌ریایی به بارندگی در هر بارش کاهش نشان می‌دهد. ظرفیت نگهداری تاج‌پوشش با استفاده از روش میانگین ۱/۸ میلی‌متر اندازه‌گیری گردید. مدل Gash، باران‌ریایی را با ۱/۱ میلی‌متر اختلاف نسبت به مقدار باران‌ریایی اندازه‌گیری شده در توده برآورد کرد. باران‌ریایی سهم قابل توجهی از بارندگی در هر بارش را در توده جنگلکاری *P.eldarica* در منطقه نیمه خشک ایران که رطوبت خاک عامل محدود کننده رشد و تولید گیاهان به حساب می‌آید، به خود اختصاص می‌دهد.