

## Long Term Sugarcane Cultivation Effects on Physical Properties of Fine Textured Soils

A. R. Barzegar<sup>1\*</sup>, Sh. Mahmoodi<sup>2</sup>, F. Hamedi<sup>3</sup> and F. Abdolvahabi<sup>4</sup>

### ABSTRACT

Long term sugarcane cultivation can alter soil physical properties. The objective of this study was to determine the effects of 38 years of sugarcane (*Saccharum officinarum* L.) cultivation on the physical properties of fine textured soils (Haplustepts and Calcustepts) in southwestern Iran. Six paired soil profiles (three pairs in each soil type) were dug along parallel transects (100 m apart) in a side-by-side comparison of areas cultivated and uncultivated with sugarcane. Composite and undisturbed soil samples from 0-30, 30-60 and 60-90 cm layers were taken for physical and chemical analysis. The results of this study revealed that long term sugarcane cultivation resulted in a higher bulk density, lower structural stability, and an increased proportion of fine pores. The decreased number of macropores together with decreased structural stability indicated the formation of a dense compacted layer at 30-60 cm in sugarcane cultivated soils.

**Keywords:** Sugarcane, Subsoil compaction, Thin section, Pore size distribution, Water characteristics curve.

### INTRODUCTION

Commercial irrigated sugarcane cultivation in Khuzestan province in southwest Iran commenced in the late 1950's. Over the last 38 years, the annual cane yield has ranged from 50 to 110 Mg ha<sup>-1</sup>. The use of heavy machinery during planting, harvesting and transporting operations in fine textured soils has led to the concern that subsoil compaction may decline long term productivity.

Hadas (1994) reviewed the theoretical analysis and experimental data on soil compaction under high axle load. He stated that subsoil compaction occurred under specific conditions, namely wet, homogenous, and deep soil under high contact pressure. Axle loads exceeding 90 kN m<sup>-3</sup> increased subsoil compaction (Salire *et al.*, 1994). Subsoil

compaction can cause serious root restriction (Tardieu, 1994; Westermann and Sojka, 1996; Håkansson *et al.*, 1996) and the loss of both transmission and water storage pores. These changes result in lower water infiltration due to the loss of transmission pores and higher soil water caused by the loss of storage pores (Soane *et al.*, 1982; Gupta *et al.*, 1987; Hadas, 1994; Lipiec *et al.*, 1998), that may consequently reduce nutrition uptake and crop yield (Hammel, 1994; Westermann and Sojka, 1996; Håkansson *et al.*, 1996; Grath and Arvidsson, 1997). Torres *et al.* (1990) reported a decrease in sugarcane rooting depth and crop yield as a result of subsoil compaction.

Different parameters and methods are used to characterize soil compactibility, such as dry bulk density, vane shear strength, pore size distribution, gas and water diffusion,

<sup>1</sup> Department of Soil Science, College of Agriculture, Shahid Chamran University, Ahwaz, Islamic Republic of Iran.

<sup>2</sup> Department of Soil Science, College of Agriculture, Tehran University, Karaj, Islamic Republic of Iran.

<sup>3</sup> Soil and Water Research Institute, Kermanshah, Islamic Republic of Iran.

<sup>4</sup> Soil and Water Research Institute, Karaj, Islamic Republic of Iran.

\* Corresponding author, e-mail: barzegar@pssci.umass.edu



and morphological analysis (Hadas, 1994). Micromorphological studies were extensively used to characterize soil deformation (Jager *et al.*, 1983; Koppi *et al.*, 1992). Jager *et al.* (1983) indicated that macropores and mesopores were less frequent in plough pans as compared with other soil layers, whereas micropores were more common. However, little information is available on the micromorphology of compacted subsoil layers, particularly under sugarcane cultivation. In this paper, changes in soil physical properties related to compaction behavior of sugarcane cultivated soils (Haft-Tapeh) were identified and compared with nearby uncultivated soils.

## METHODS AND MATERIALS

### Soil Sampling and Analysis

This research was conducted on a commercial sugarcane plantation in Haft-Tapeh (32° 04' N, 48° 21' E), southwestern Iran. Prior to the planting of sugarcane in 1959, the site was under a wheat (*Triticum* sp.) base system. The area has a unimodal rainfall with a dry season from May to November and a rainy season from November to April. The mean annual rainfall of the area studied is 270 mm. Evaporation (Class A open pan) is about 2700 mm yr<sup>-1</sup>. The mean annual temperature is 31.6° C. The soils are pedologically young and have been developed on transported sediments.

Six paired soil profiles (three in each soil

type) were dug along parallel transects (100 m apart) in a side-by-side comparison of areas cultivated and uncultivated with sugarcane. The transects crossed two soil families: fine, mixed, hyperthermic Haplustepts and fine loamy, mixed, hyperthermic Calcustepts (Table 1). The uncultivated soils exist between the main drain and access road. These soils have not been cultivated for about four decades.

Twenty-four undisturbed soil cores were taken from 30-60 and 60-90 cm layers of soil profiles at farms 101, 102 and 103 (hereafter referred to as I, II and III, respectively) of the Haft-Tapeh sugarcane complex. Undisturbed soil samples were not taken from 0-30 cm because of heavy disturbance of this layer during plantation. The water content at sampling time (October, 1998) varied from 60% to 80% of the field capacity. The undisturbed soil samples were used to determine dry bulk density, water characteristic curves and structural stability. The water characteristics curves were determined on a pressure membrane apparatus, using three intact cores taken from each soil layer. Three measurements were made in each intact core. The pressures used were 10, 30, 50, 100, 1000 and 1500 kPa. The pore size distribution was calculated using the water characteristics curves (Jury and Horton, 2004) and the structural stability of 2-4 mm aggregates was determined by a method outlined in Barzegar *et al.* (1996). The aggregates were oven dried for 72 hours at 50° C. Five grams of the aggregates were then shaken with distilled water (25 mL) for

**Table 1.** Classification of soils (Soil Survey Staff, 1998; FAO/UNESCO, 1998).

Soil type	Family	Subgroup	FAO/UNESCO
I <sub>SC</sub> <sup>a</sup>	Fine, mixed, hyperthermic	Calcustepts	Calcic Cambisol
I <sub>UC</sub> <sup>b</sup>	Fine loamy, mixed, hyperthermic	Haplustepts	Utric Cambisol
II <sub>SC</sub>	Fine loamy, mixed, hyperthermic	Haplustepts	Utric Cambisol
II <sub>UC</sub>	Fine loamy, mixed, hyperthermic	Calcustepts	Calcic Cambisol
III <sub>SC</sub>	Fine, mixed, hyperthermic	Haplustepts	Calcic Cambisol
III <sub>UC</sub>	Fine loamy, mixed, hyperthermic	Haplustepts	Calcic Cambisol

<sup>a</sup> SC, sugarcane cultivated soils; <sup>b</sup> UC, uncultivated soils

5 minutes on an end-over-end shaker. The suspension was passed through sieves of 250 $\mu$ m with three additional washes of 50 mL of distilled water. The particles remaining on the sieves were oven dried at 105° C and weighed. The sand content of the macro aggregates (> 250 $\mu$ m) was also determined.

Composite soil samples from 0-30, 30-60 and 60-90 cm of each soil profile were used to measure chemical properties and particle size distribution. Organic matter content was determined by the dichromate oxidation method (Nelson and Sommers, 1982). The cation exchange capacity (CEC) was measured by a Na-saturation procedure (Rhoades, 1982a). Electrical conductivity of the saturation extract (EC<sub>e</sub>) and the pH of the saturated paste of soils were measured (Rhoades, 1982b). Particle size distribution was determined by the pipette method (Gee and Bauder, 1986).

### Thin Sections

Thirty two samples were taken from 30-60 and 60-90 cm of paired profiles (four samples from each). Thin sections for study under a polarizing-light microscope were then prepared (Murphy, 1986). Undisturbed samples were impregnated with a polyester resin-acetone mixture under a partial vacuum. The resin permeated into the soil and set hard in about 48 hours. The surface was removed with a hard and soft grinder to expose a fresh, flat surface of resin-impregnated mixture. The samples were then sliced, smoothed and mounted on glass. The mounted blocks were also thinned and smoothed with a diamond-edge saw, lapping machine and corundum powder to reach a thickness of 25 to 30 $\mu$ m. The central faces of the 100mm $\times$ 100mm area were photographed using a 35-mm camera fitted with a macrolens. Image analysis was carried out using a special program (Image tool) and various soil structural characteristics were evaluated. Although image analysis was used to analyze a two-dimensional picture, this process provides useful information

about the complexity of pore patterns in soils, which could not be obtained with other methods, such as mercury intrusion, water retention, and nitrogen sorption (Bullock *et al.*, 1985). Thin sections were described using the manual written by Bullock *et al.* (1985).

### Statistical Analysis

Student's *t* test was used for a paired comparison of cultivated and uncultivated soil samples. Least Significant Difference (LSD) values were calculated to compare pairs of mean values at either 5% or 1% levels of probability. The statistical analysis was performed using the SAS statistical package (SAS Institute, 1988).

## RESULTS AND DISCUSSION

### Physical and Chemical Properties of Soils

The physical and chemical characteristics of the soils are presented in Table 2. The soils under sugarcane (except the surface layer of soil II) had a higher clay ( $P < 0.05$ ) and lower sand ( $P < 0.05$ ) content compared to uncultivated soils. The differences in particle size could be due to deep tillage and subsoiling use for sugarcane cultivation. The organic carbon content of 0-30 cm of sugarcane cultivated soils was more than that of uncultivated soils ( $P < 0.01$ ); for other layers the difference was not significant. The cation exchange capacity (CEC) of sugarcane cultivated soils was higher at 0-30 and 30-60 cm than in uncultivated soils ( $P < 0.01$ ) due to differences in both organic matter and clay contents. Sugarcane cultivated soils had lower electrical conductivity and pH values as compared to uncultivated soils ( $P < 0.01$ ). Excess leaching of salts in sugarcane cultivated soils resulted in lower soil electrical conductivity. In salt-affected soils, leaching excess salts increases the pH of soils (Barzegar, 2000).



The bulk density of the 30-60 cm layer of sugarcane cultivated soils was higher than that of uncultivated soils ( $P < 0.01$ ). No difference was found among the bulk densities of the 60-90 cm layers. Heavy duty machinery and increasing numbers of passage of agricultural machinery over the long term result in subsoil compaction (Soane *et al.*, 1982; Håkansson *et al.*, 1996). However, the bulk density of the 0-30 cm layer of sugarcane cultivated soils was lower than that of uncultivated soils ( $P < 0.05$ ). This might be

due to either cane residuals left after harvesting or soil disturbance after cultivation. Barzegar *et al.* (2000) investigated the influence of different amounts of sugarcane residue and soil water content on soil compactibility at different energy input levels. They indicated that sugarcane residue was effective in reducing soil compactibility at various soil water contents and compaction levels. These results differ from those reported by Hartemink (1998). He compared the physical and chemical properties of soils

**Table 2.** Physical and chemical characteristics of selected soil layers.

Soil type	Depth (cm)	Sand	Silt	Clay	O.C <sup>a</sup>	BD <sup>b</sup> (Mg m <sup>-3</sup> )	CEC <sup>c</sup> (Cmol kg <sup>-1</sup> )	ECe <sup>d</sup> (dS m <sup>-1</sup> )	pH
I <sub>SC</sub> <sup>e</sup>	0-30	150	450	400	10.0	1.35	16.5	0.8	8.0
I <sub>UC</sub> <sup>f</sup>	0-30	230	460	310	5.0	1.41	12.0	16.1	7.2
SL <sup>g</sup>		*	ns	*	**	*	**	**	**
I <sub>SC</sub>	30-60	170	430	400	5.0	1.60	16.0	0.7	8.1
I <sub>UC</sub>	30-60	285	443	272	3.7	1.42	12.5	5.7	7.6
SL		**	ns	**	ns	**	**	**	**
I <sub>SC</sub>	60-90	160	490	350	3.5	1.50	14.0	0.5	8.0
I <sub>UC</sub>	60-90	270	455	275	2.0	1.50	12.3	7.1	7.6
SL		**	ns	*	ns	ns	ns	**	*
II <sub>SC</sub>	0-30	240	490	270	8.2	1.36	12.0	0.6	8.0
II <sub>UC</sub>	0-30	210	470	320	5.0	1.40	12.3	10.1	7.4
SL		ns	ns	ns	*	ns	ns	**	**
II <sub>SC</sub>	30-60	170	500	330	4.0	1.55	14.0	0.5	8.1
II <sub>UC</sub>	30-60	250	440	310	2.5	1.40	13.5	6.5	7.6
SL		*	*	ns	ns	**	ns	**	**
II <sub>SC</sub>	60-90	200	450	350	3.8	1.45	14.0	0.5	8.1
II <sub>UC</sub>	60-90	350	400	250	1.0	1.50	11.0	6.0	7.6
SL		*	*	**	*	ns	*	**	**
III <sub>SC</sub>	0-30	240	430	330	7.1	1.35	14.2	0.7	8.0
III <sub>UC</sub>	0-30	220	475	305	5.4	1.35	13.0	27.5	7.3
SL		ns	ns	ns	ns	ns	ns	**	**
III <sub>SC</sub>	30-60	120	485	395	4.5	1.60	16.0	0.5	8.1
III <sub>UC</sub>	30-60	320	509	171	3.4	1.42	9.5	13.8	7.6
SL		**	ns	**	ns	**	**	**	**
III <sub>SC</sub>	60-90	50	590	360	4.0	1.51	15.2	0.5	8.1
III <sub>UC</sub>	60-90	350	400	250	3.0	1.42	10.5	13.0	7.8
SL		**	**	**	ns	ns	**	**	ns

<sup>a</sup> O.C, organic carbon; <sup>b</sup>BD, bulk density; <sup>c</sup>CEC, cation exchange capacity; <sup>d</sup> ECe, electrical conductivity of saturated extract; <sup>e</sup>SC, sugarcane cultivated soil; <sup>f</sup> UC, uncultivated soils; <sup>g</sup> SL, significant level, ns not significant, \* significant at 5%, and \*\* significant at 1%.

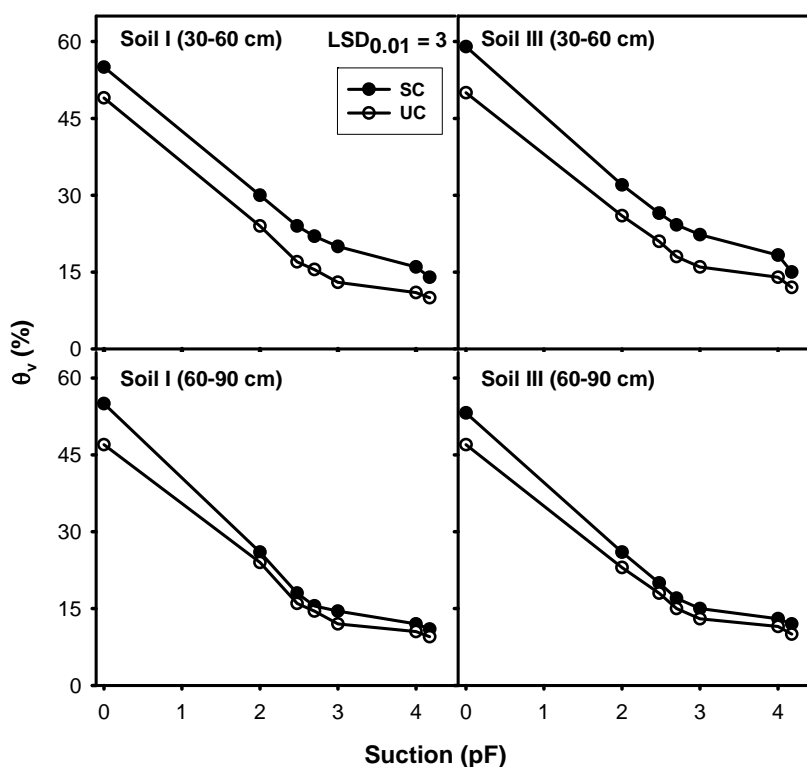
(Eutric and Mollic Fluvisols and Eutric Vertisols) with an average annual rainfall of 2000 mm under sugarcane and grassland, and indicated no significant differences between either the bulk density or infiltration rate of grassland and sugarcane rows. However, the interrow had a significantly higher bulk density and lower infiltration rate due to wheeled traffic. He also reported a decline in soil pH from 6.5 to 5.8 in sugarcane cultivated soils.

### Pore Size Distribution

Figure 1 shows the water characteristics curves of soils. Water content at saturation of soils under sugarcane was about 10% higher than that of uncultivated soils. The decrease in soil matric potential from 0 (saturation) to -33 kPa (field capacity) was very gradual in the 30-60 cm layer of uncultivated soils compared to sugarcane culti-

vated soils, indicating the presence of large voids with different diameters. The differences were small in the 60-90 cm layer, indicating fewer effects of compaction on this layer. The reductions in the wet range (from 0 to -33 kPa) were similar in both sugarcane cultivated and uncultivated soils although the absolute values of water content at given water potential were different.

The pore size distribution calculated using water characteristics curves is illustrated in Figure 2. Results indicated a lower percentage of macropores (>9.3  $\mu\text{m}$  diameter) in sugarcane cultivated soils. Pores with diameter of 9.3 to 29  $\mu\text{m}$  in uncultivated soils were greater than sugarcane cultivated soils. The reduction in macroporosity suggests a reduction in aeration and water use efficiency. Few to common fine redoximorphic features in the 30-60 cm layer of sugarcane cultivated soils (revealed by studying the thin sections; Figure not shown) was also the result of different hydraulic conductivities



**Figure 1.** Water characteristics curves of 30-60 and 60-90 cm layers of sugarcane cultivated (SC) and uncultivated (UC) soils.



caused by changes in the void distribution pattern. These features have not been observed on similar layers of uncultivated soils. Kooistra *et al.* (1992) investigated the influence of deep tillage and compaction on soil structure and reported similar results.

The proportion of different pore shapes is presented in Table 3. The percentages of

shapes (Table 3) combined with pore size distribution (Figure 2) showed a decline in macropores in the subsoil (30-60 cm) layers of sugarcane cultivated soils ( $P < 0.01$ ).

### Structural Stability

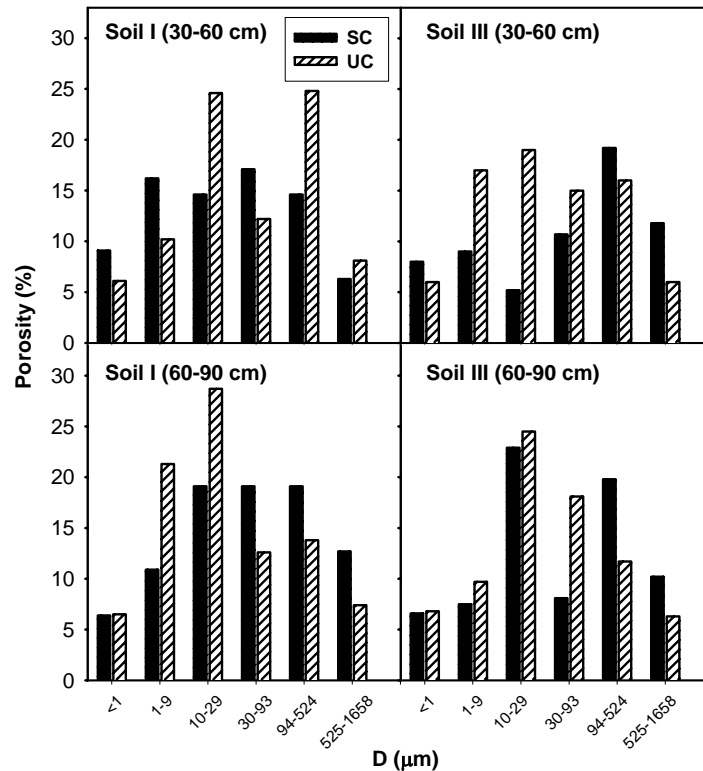


Figure 2. Pore size distribution of soils calculated from the water characteristics curves.

channel, chamber and vughy pores were generally greater in the 30-60 cm layer than in the 60-90 cm layer, probably due to the higher biological activity. The same trend was observed for sugarcane cultivated soils. Overall, the number of channel, chamber and Vughy pores were higher in sugarcane cultivated soils as compared to uncultivated soils ( $P < 0.01$ ). The percentage of vesicle and packing voids were similar in both sugarcane cultivated and uncultivated soils. The results of quantitative measurements of pore

The macroaggregate ( $>250 \mu\text{m}$ ) stability of soil layers is given in Table 4. Generally, the proportion of macroaggregates in the 30-60 cm layer of sugarcane cultivated soils was relatively lower compared to uncultivated soils. However, in the topsoil of sugarcane cultivated soils, the water stable macroaggregate stability was greater than in uncultivated soils. This can be due to both the higher clay and organic matter content of these soils.

**Table 3.** The percentage of different types of pores calculated using the thin sections.

Soil type	Depth (cm)	Channels	Chambers	Vughs	Planes	Vesicles	Packing voids
I <sub>SC</sub> <sup>a</sup>	30-60	42	35	10	5	3	5
I <sub>UC</sub> <sup>b</sup>	30-60	26	25	25	14	5	5
SL <sup>c</sup>		**	**	**	**	ns	ns
I <sub>SC</sub>	60-90	42	40	5	6	3	4
I <sub>UC</sub>	60-90	25	30	20	15	5	5
SL		**	**	**	**	ns	ns
III <sub>SC</sub>	30-60	35	35	15	5	5	5
III <sub>UC</sub>	30-60	25	20	35	15	2	3
SL		**	**	**	**	ns	ns
III <sub>SC</sub>	60-90	50	10	10	20	5	5
III <sub>UC</sub>	60-90	25	20	30	15	5	5
SL		**	**	**	**	ns	ns

<sup>a</sup> SC, sugarcane cultivated soil; <sup>b</sup> UC, uncultivated soils; <sup>c</sup> SL, significant level; <sup>ns</sup> not significant; \* significant at 5% and \*\* significant at 1%

All the attributes measured, including bulk density, water characteristics curve and macroaggregate stability revealed that soil structural stability in the 30-60 cm layer of sugarcane cultivated soils was decreased. Cumulative heavy traffic for planting, harvesting and transporting sugarcane in the last 38 years has led to subsoil compaction at a depth of 30-60cm. There are a number of possible reasons for the formation of compacted subsoil at the Haft-Tapeh. The harvesting season of sugarcane in Khuzestan province lasts from September to March. This period is the rainy season of this area. Heavy machinery traffic on sugarcane cultivated soils with the optimum water content for compaction will result in soil compaction. Barzegar *et al.* (2000) investigated the compactibility of sugarcane cultivated soils and reported that soil water content near the plastic limit results in a favorable conditions for soil compaction. The compactibility of soils is also dependent on the weight, contact area and speed of agricultural machinery (Soane *et al.*, 1982; Soane and van Ouwerkerk, 1994). The weight of the agricultural machinery used for sugarcane production at

Haft-Tapeh including transporters, trucks and vanguards are 18.3, 20, 43.5 Mg, respectively. The contact areas of these are 35.61, 7.46 and 12.93 m<sup>2</sup>, respectively. The heavy weight and low contact area of vanguards would be one the main reason for soil compaction. The other reason for the formation of a compacted subsurface layer in sugarcane cultivated soils might be the low organic matter content of these soils. Most of the sugarcane crops are harvested after pre-harvesting burning. This will reduce the amount of residue left on soils after harvesting. Compaction of the plough layer of sugarcane cultivated soils can be alleviated by restricting heavy machinery traffic at a water content less than 0.8 of the plastic limit, and incorporating sugarcane residue at a rate of 27 Mg ha<sup>-1</sup> (Barzegar *et al.*, 2000). Crop residues with very low decomposition rates not only increase the C:N ratio (Oades, 1984) but also may interfere with common agricultural practices and the movement of irrigation water in furrows. Barzegar *et al.* (2002) indicated that the C:N of composting bagasse (C:N of 25) was almost one fifth of sugarcane residue (C:N of 135). A number

**Table 4.** Water stable macroaggregates (>250 µm; MA) of sugarcane cultivated (SC) and uncultivated (UC) soils.

Soil type	Depth (cm)	MA <sub>SC</sub> <sup>a</sup>	MA <sub>UC</sub> <sup>b</sup>	<sup>c</sup> Significance level
		%		
I	0-30	40	32	**
	30-60	28	29	ns
	60-90	32	25	**
II	0-30	29	35	ns
	30-60	25	30	*
	60-90	30	25	ns
III	0-30	40	34	*
	30-60	25	29	*
	60-90	30	20	**

<sup>a</sup> SC, sugarcane cultivated soil; <sup>b</sup> UC, uncultivated soils; <sup>c</sup> ns, not significant; \* significant at 5% and \*\* significant at 1%

of studies have illustrated that organic matter increased soil structural stability (Kay, 1990; Barzegar *et al.*, 2002), water holding capacity at higher soil water potential (i.e. between 0 to -100 kPa), soil porosity and the infiltration rate (Barzegar *et al.*, 2002) and decreased soil compactibility (Barzegar *et al.*, 2000). Viator *et al.* (2002) recommended subsoil rather than within-row application of compost as the preferred practice for sugarcane production in heavy textured soils.

### CONCLUSION

Results indicated that long term sugarcane cultivation altered soil physical properties. Aggregate stability and macropore proportions decreased and bulk density increased at a depth of 30-60 cm of sugarcane cultivated soils. An aquic condition was likely developed at the compacted subsurface layer of sugarcane cultivated soils as revealed by redoximorphic features.

### ACKNOWLEDGEMENTS

We are grateful to Professor Alfred E. Hartemink of the International Soil Reference and Information Center (ISRIC), The Neth-

erlands, and three anonymous reviewers for their comments.

### REFERENCES

1. Barzegar, A. R. 2000. *Salt-affected Soils: Diagnosis and Productivity*. Shahid Chamran University Press, Ahwaz, Iran. (In Persian).
2. Barzegar, A. R., Asoodar, M. A. and Ansari, M. 2000. Effectiveness of Sugarcane Residue Incorporation at Different Water Contents and the Proctor Compaction Loads in Reducing Soil Compactibility. *Soil and Tillage Res.*, **57**: 167-172.
3. Barzegar, A. R., Nelson P. N., Oades J. M. and Rengasamy P. 1996. Soil Structure Degradation and Mellowing of Compacted Soils by Saline-sodic Solutions. *Soil Sci. Soc. Am. J.*, **60**: 583-588.
4. Barzegar, A. R., Yousefi A. and Daryashenas A. 2002. The Effect of Addition of Different Amounts and Types of Organic Materials on Soil Physical Properties and Yield of Wheat. *Plant Soil.*, **247**: 295-301.
5. Bullock, P., Fedoroff, N., Jongerius, A., Stoops, G., Tursina, T. and Babel, U. 1985. *Handbook for Soil Thin Section Description*. Waine Research Publications. Wolverhampton, UK.
6. FAO-UNESCO. 1998. *Soil Map of the World. Revised Legend Required as International Soil Reference and Information Cen-*



- ter. Technical Paper No. 20. Wageningen, ISRIC.
7. Gee, G. W. and Bauder J. W. 1986. Particle Size Analysis. In: *Methods of Soil Analysis, Part 1, Physical Properties*. 2<sup>nd</sup> Ed., (Ed.): Klute, A., ASA and SSSA, Madison, WI, pp. 377-381.
  8. Grath, T. and Arvidsson, J. 1997. Effect of Soil Compaction on Plant Nutrition Uptake and Growth of Peas and Barley on a Sandy Loam. *Swedish J. Agric. Res.*, **27**: 95-104.
  9. Gupta, S. C., Schneider, E. C., Larson, W. E. and Hadas, A. 1987. Influence of Corn Residue on Compression and Compaction Behavior of Soils. *Soil Sci. Soc. Am. J.*, **51**: 207-212.
  10. Hadas, A. 1994. Soil Compaction Caused by High axle Loads: Review of Concepts and Experimental Data. *Soil Tillage Res.*, **29**: 253-276.
  11. Håkansson, I., Grath, T. and Olsen, H. J. 1996. Influence of Machinery Traffic in Swedish Farm Fields on Penetration Resistance in the Subsoil. *Swedish J. Agric. Res.*, **26**: 181-187.
  12. Hammel, J. E. 1994. Effect of High-axle Load Traffic on Subsoil Physical Properties and Crop Yield in the Pacific Northwest USA. *Soil Tillage Res.*, **29**: 193-203.
  13. Hartemink, A. E. 1998. Soil Chemical and Physical Properties as Indicators of Sustainable Land Management under Sugar Cane in Papua New Guinea. *Geoderma*, **85**: 283-306.
  14. Jager, A., Bouma, J. and Bisdom, E. D. A. 1983. The Characterization of Microporosity in a Ploughpan by Submicroscopic and Quatimet Techniques. *Geoderma*, **30**: 277-283.
  15. Jury, W. A. and Horton, R. 2004. *Soil Physics*. 6th Edition. John Wiley & Sons, NY.
  16. Kay, B. D. 1990. Rates of Changes of Soil Structure under Different Cropping Systems. *Adv. Soil Sci.*, **12**: 1-52.
  17. Kooistra, M. J., Schoonderbeek, D., Boone, F. R., Veen, B. W. and van Noordwijk, M. 1992. Root-Soil Contact of Maize as Measured by a Thin-Section Technique. II. Effects of Soil Compaction. *Plant Soil*, **139**: 119-129.
  18. Koppi, A. J., Douglas, J. T. and Moran, C. T. 1992. An Image Analysis Evaluation of Soil Compaction in Grassland. *J. Soil Sci.*, **43**: 15-25.
  19. Lipiec, J., Hatano, R. and Sowiska-Jurkiewicz, A. 1998. The Fractal Dimensions of Pore Distribution Pattern in Various Compacted Soils. *Soil Tillage Res.*, **47**: 61-66.
  20. Murphy, C. P. 1986. *Thin Section Preparation of Soils and Sediments*. Rothamsted Experimental Station. ABA Publ., UK.
  21. Nelson, D. W. and Sommers, L. E. 1982. Total Carbon, Organic Carbon, and Organic Matter. In: *Methods of Soil Analysis, Part 2, Chemical and Microbiological Properties*. 2nd edition. (Eds.): Page, A. L., Miller, R. H. and Keeney, D. R., Agron. Monogr. 9. ASA and SSSA, Madison, WI, pp. 539-573.
  22. Oades, J. M. 1984. Soil Organic Matter and Structural Stability: Mechanisms and Implications for Management. *Plant Soil*, **76**: 319-337.
  23. Rhoades, J. D. 1982a. Cation Exchange Capacity. In: *Methods of Soil Analysis, Part 2, Chemical and Microbiological Properties*. 2nd edition. (Eds.): Page, A. L., Miller, R. H. and Keeney, D. R., Agron. Monogr. 9. ASA and SSSA, Madison, WI, pp. 149-157.
  24. Rhoades, J. D. 1982b. Soluble Salts. In: *Methods of Soil Analysis, Part 2, Chemical and Microbiological Properties*. 2nd edition. (Eds.): Page, A. L., Miller, R. H. and Keeney, D. R., Agron. Monogr. 9. ASA and SSSA, Madison, WI, pp. 167-178.
  25. Salire, E. V., Hammel, J. E. and Hardcastle, J. H. 1994. Compression of Intact Subsoils under Short-duration Loading. *Soil Tillage Res.*, **31**: 253-248.
  26. SAS Institute. 1988. *SAS User's Guide: Statistical*. 6<sup>th</sup> edition. SAS Inst., Cary, NC.
  27. Soane, B. D. and van Ouwkerk, C. 1994. *Soil Compaction in Crop Production*. Elsevier, Bonn, Germany.
  28. Soane, B. D., Dikson, J. W. and Campbell, D. J. 1982. Compaction by Agricultural Vehicle: A Review. III. Incidence and Control of Compaction in Crop Production. *Soil Tillage Res.*, **2**: 30-36.
  29. Soil Survey Staff. 1998. *Key to Soil Taxonomy*. 7<sup>th</sup> Edition. Blacking Pocahontas Press Inc., USDA.
  30. Tardieu, F. 1994. Growth and Functioning of Roots and Root Systems Subjected to Soil Compaction: Toward a System with Multiple Signaling. *Soil Tillage Res.*, **29**: 217-243.
  31. Torres, J. S., Yang, S. J. and Villegas, F. 1990. Soil Compaction and Cane Stalk Damage by Semi-mechanized Harvesting Systems in the Wet Season. *Sugar Cane*, **5**: 12-16.



32. Viator, R. P., Kovar, J. L. and Hallmark, W. B. 2002. Gypsum and Compost Effects on Sugarcane Root Growth, Yield, and Plant Nutrients. *Agron. J.*, **94**: 1332-1336.
33. Westermann, D. T. and Sojka, R. E. 1996. Tillage and Nitrogen Placement Effects on Nutrient Uptake by Potato. *Soil Sci. Soc. Am. J.*, **60**: 1448-1453.

## اثرات طولانی مدت کشت نیشکر بر خصوصیات فیزیکی خاکهای ریزبافت

ع. ر. برزگر، ش. محمودی، ف. حامدی و ف. عبدالوهابی

### چکیده

کشت طولانی مدت نیشکر می‌تواند باعث تغییرات عمده‌ای در خصوصیات فیزیکی خاکها شود. هدف از انجام این تحقیق بررسی اثر 38 سال کشت نیشکر (*Saccharum officinarum*) بر خصوصیات فیزیکی خاکهای ریزبافت (هپل یوستپت و کلسی یوستپت) در جنوب غرب ایران می‌باشد. شش پروفیل دوتایی (سه زوج پروفیل در هر تیپ خاک) در مزارع کشت نیشکر و در خاکهای کشت نشده، به موازات هم در طول ترانسکت‌ه‌ای به فاصله 100 متر حف گردید. از اعماق 0-30، 30-60 و 60-90 سانتیمتری این پروفیل‌ها نمونه‌های مخلوط و دست‌نخورده جهت آزمایشات فیزیکی، شیمیایی و میکرومرفولوژیکی تهیه گردید. نتایج این تحقیق نشان داد که کشت طولانی مدت نیشکر باعث افزایش جرم مخصوص ظاهری، کاهش پایداری ساختمان و افزایش خلل و فرج ریز خاک گردیده است. کاهش خلل و فرج درشت به‌مراه کاهش پایداری ساختمان خاک نشان از تشکیل لایه فشرده و متراکم در عمق 30-60 سانتیمتری خاکهای تحت کشت نیشکر دارد.